

Motivation

1. Flood risk analyses that account for:
 - Climate Changes
 - River Channel-Floodplain Geomorphology
2. Prediction in gaged & un-gaged basins
3. Watershed perspective for planning & policy
4. Hydrologic resilience of river basins

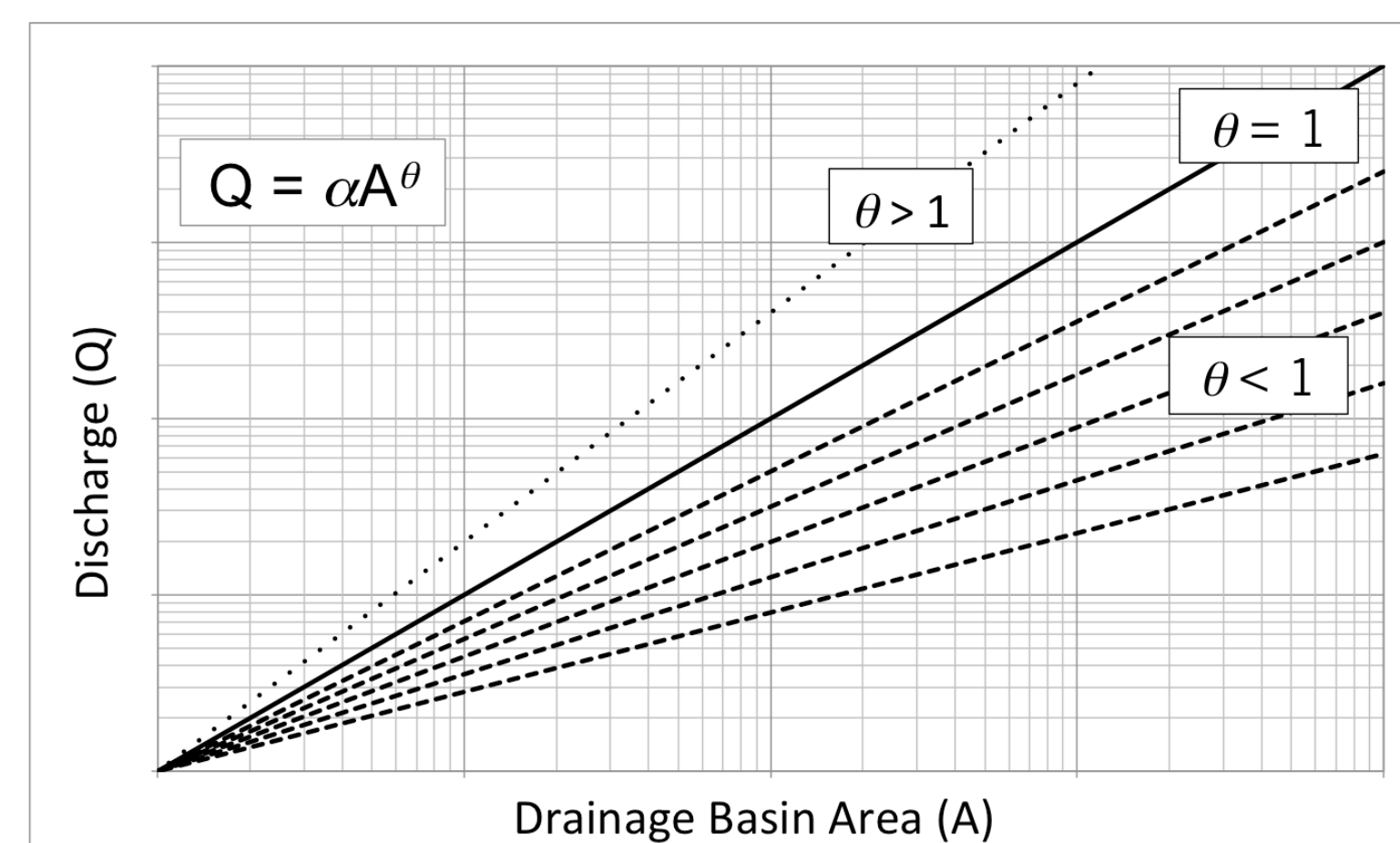
Hydrologic Resilience

- Ability of a river network to accommodate hydro-climatic shifts without significant changes to channel-floodplain morphology or functioning

Study Approach

- Discharge-Basin Area relationships widely used in regional flood frequency regression models
- Examine relationships not just for prediction, but *understanding* of physical mechanisms controlling flood magnitude

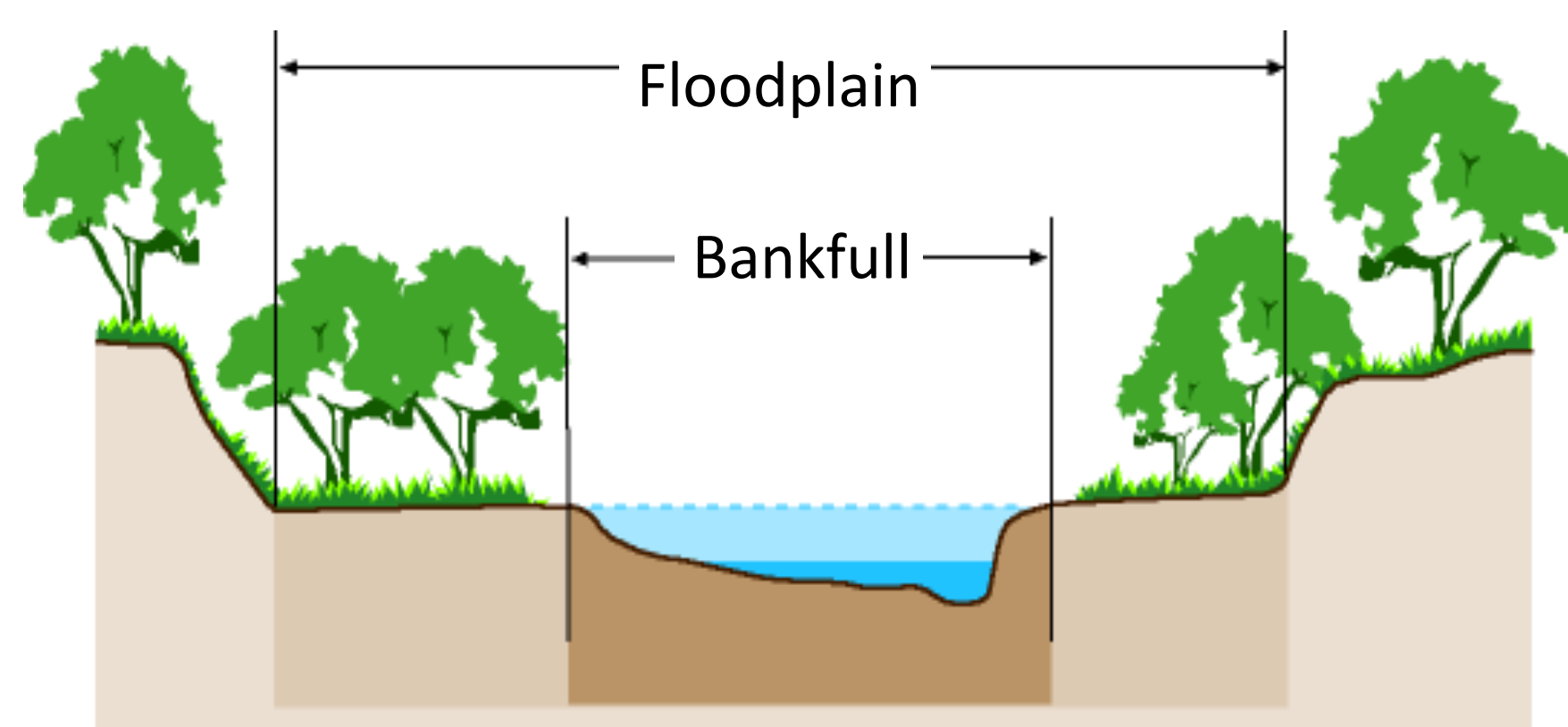
Spatial Scaling of Floods



For scaling of **floods**: Q can be peak discharge of event or flood frequency quantile estimate. Most floods have scaling exponents (θ) less than 1.

- Why is θ less than 1?
- What causes variation in θ ?

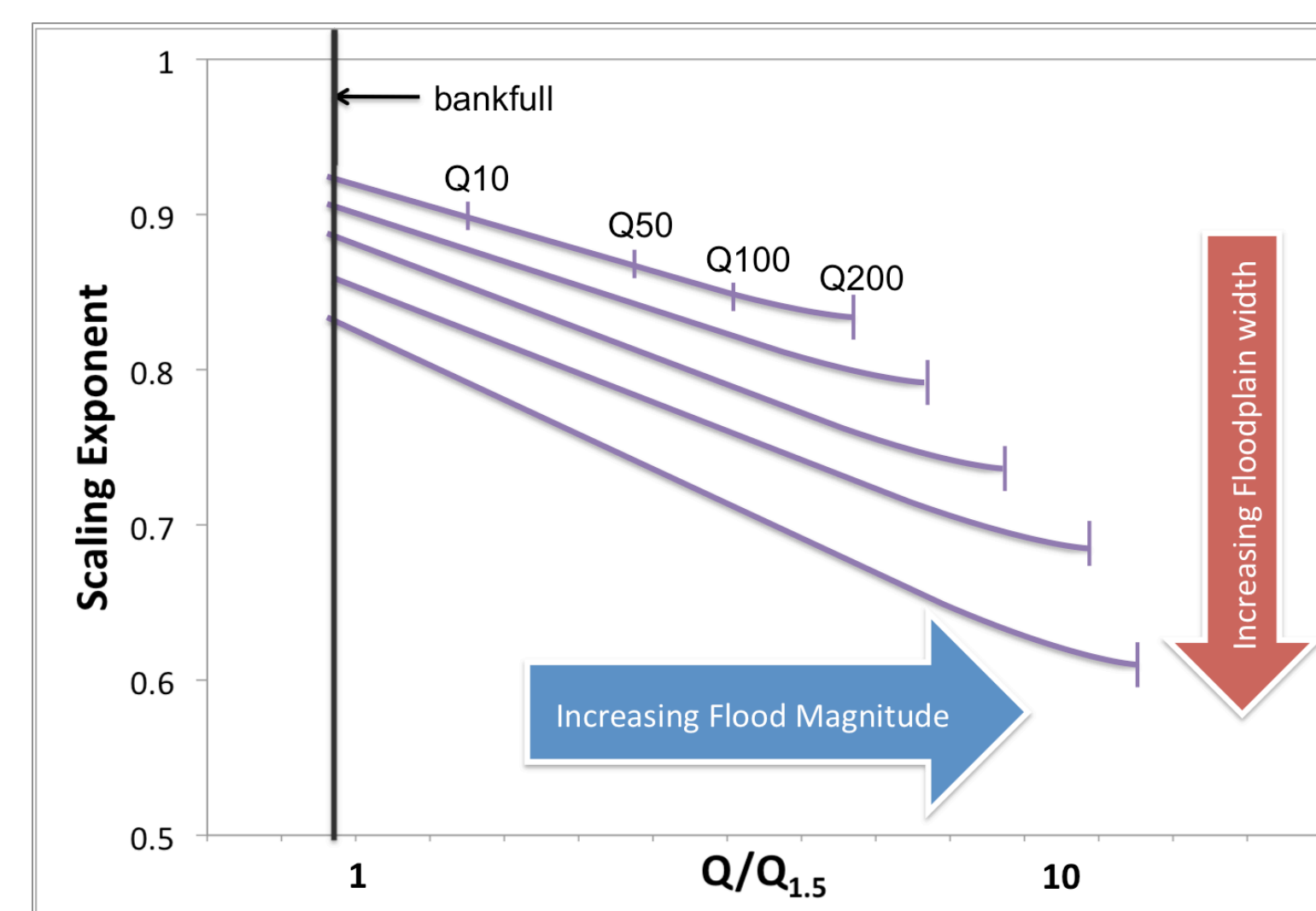
Channel-floodplain morphology



<http://ag.arizona.edu/watershedsteward/resources/>

- Bankfull flow** transports sediment, maintains channel shape and integrates the channel network
- expect scaling exponents closer to 1
- Floods** flow onto floodplain and dissipate flood peaks
- expect lower scaling exponents with larger floods

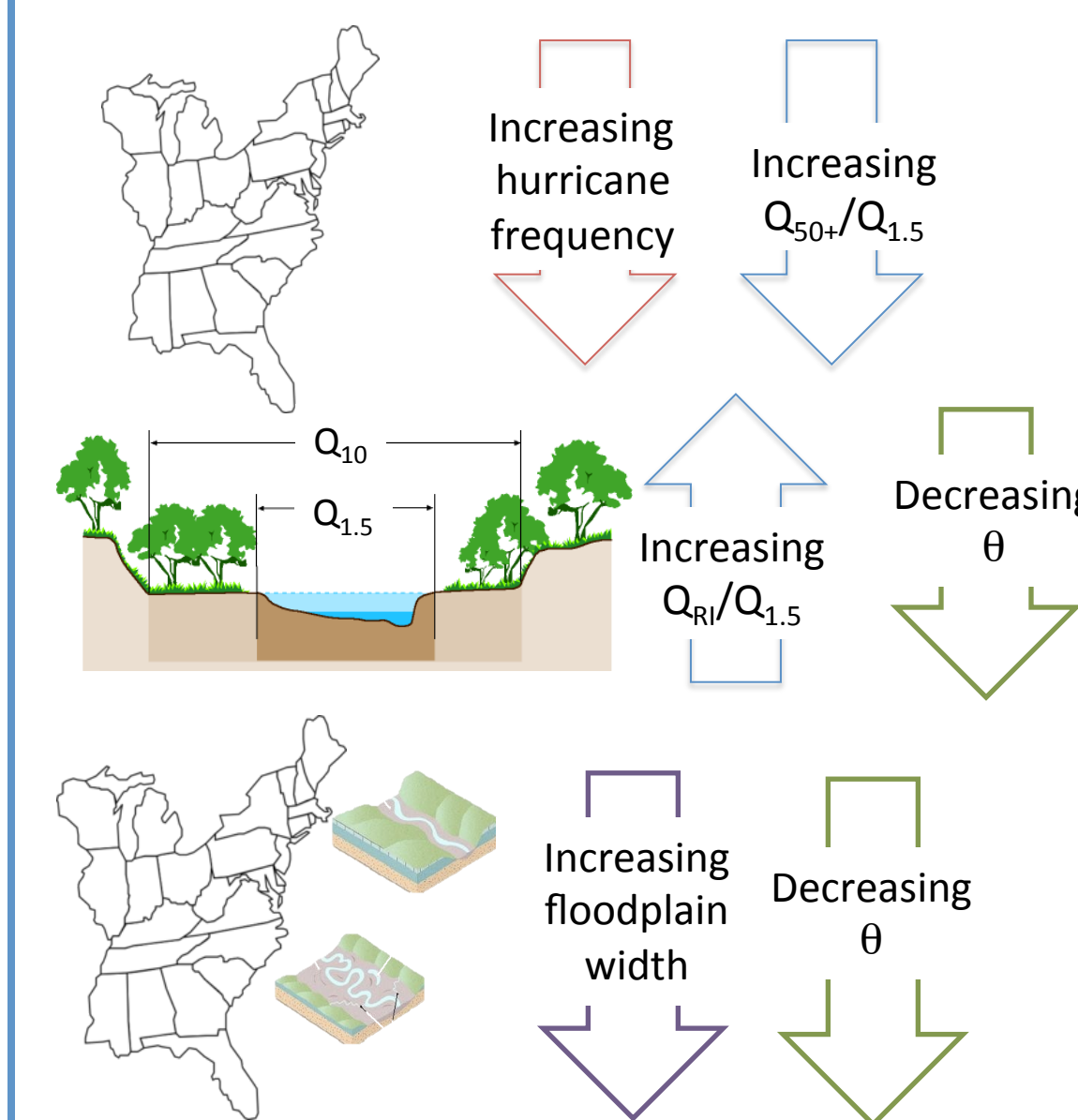
Conceptual Model



Scaling exponents should **decrease** with:

- **increasing** flood magnitude: greater peak dissipation onto floodplain
- **increasing** floodplain width: greater capacity to dissipate flood peaks

Hypotheses

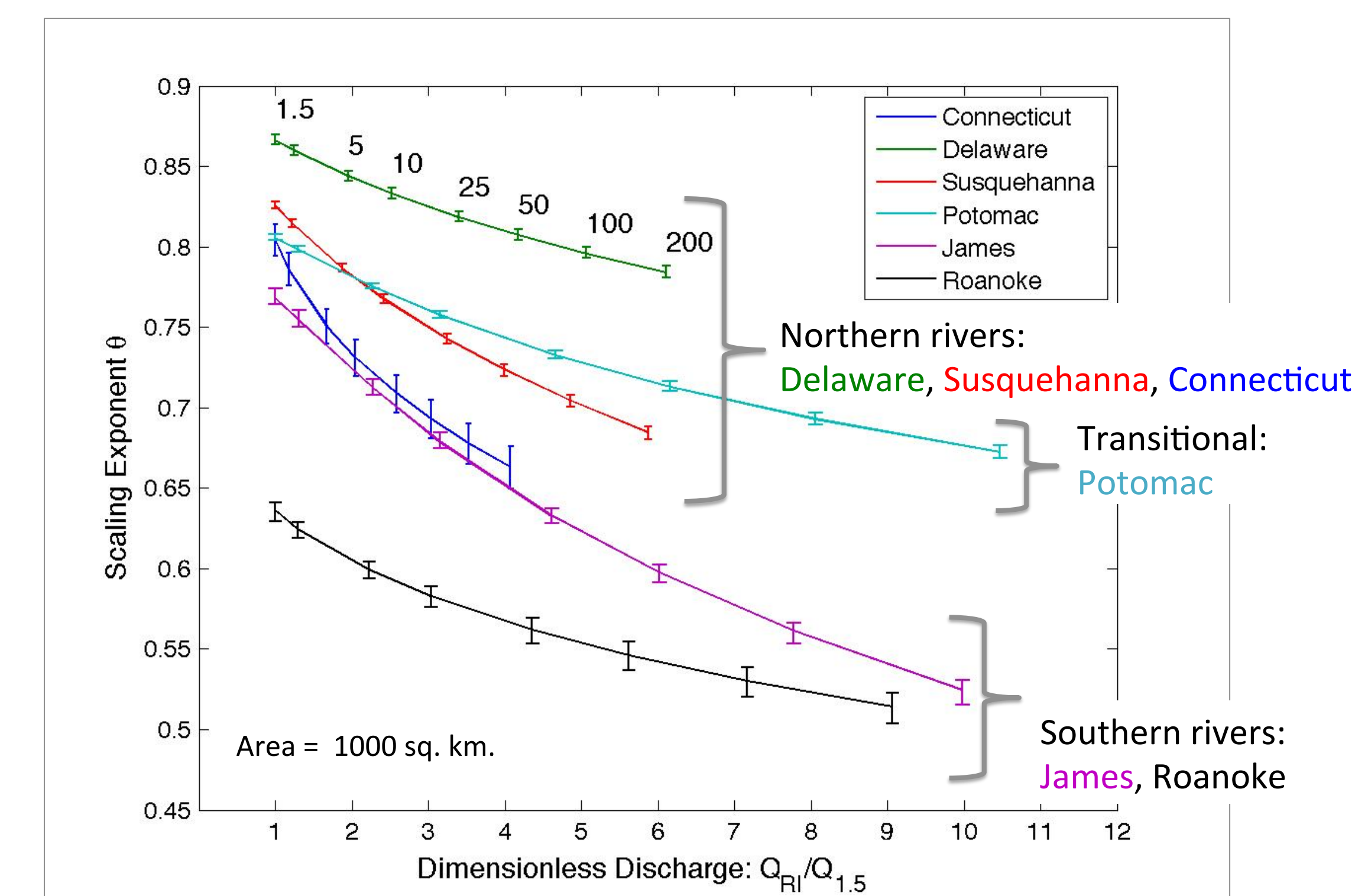


H1: largest dimensionless discharges (for $R1 > 50$) are in southeastern rivers

H2: lower scaling exponents for larger floods in all rivers

H3: Rivers with wider floodplains will have the lowest scaling exponents

Results



Scaling exponents:

- **Decrease** as dimensionless discharge increases in all river basins
- Are **higher** in northern river basins: Connecticut, Delaware, Susquehanna
- Are **lower** in southern river basins: James, Roanoke

Potomac River basin is **transitional**

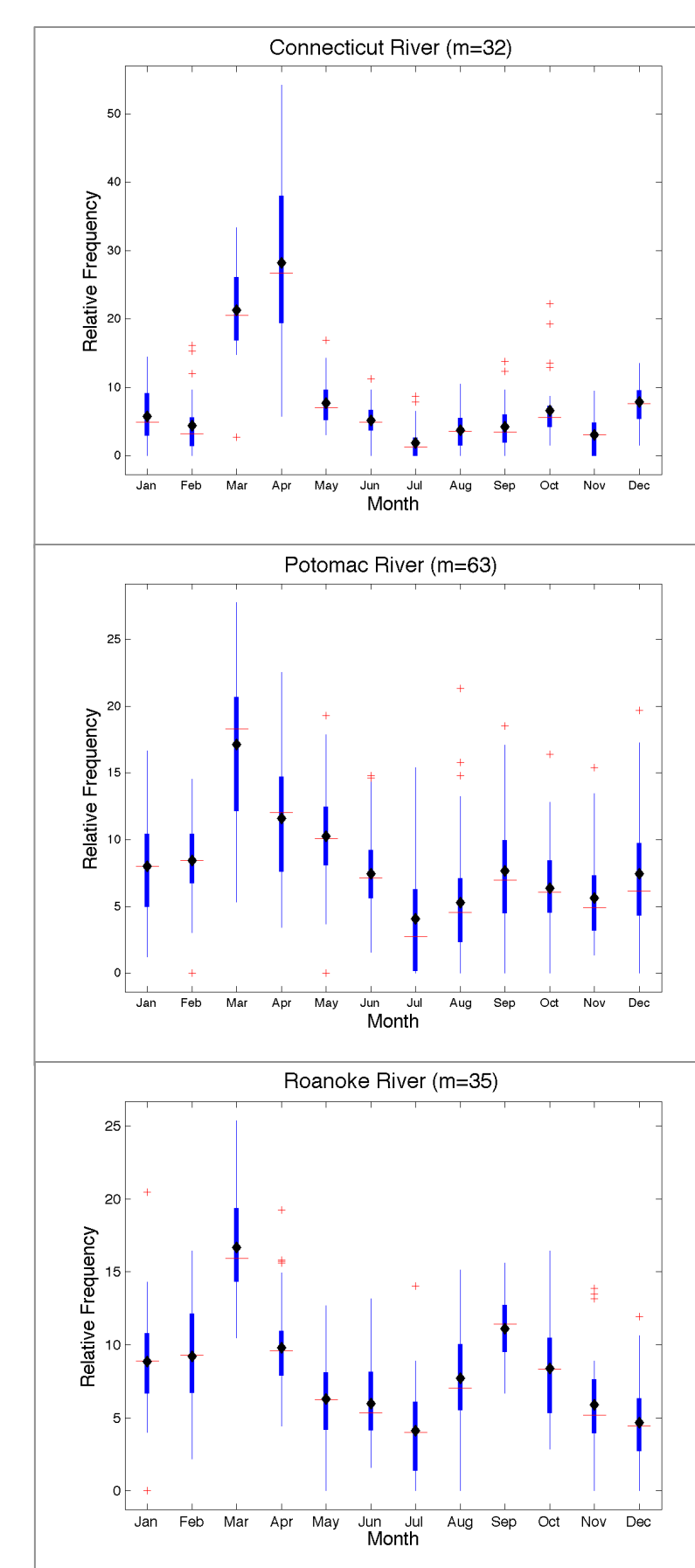
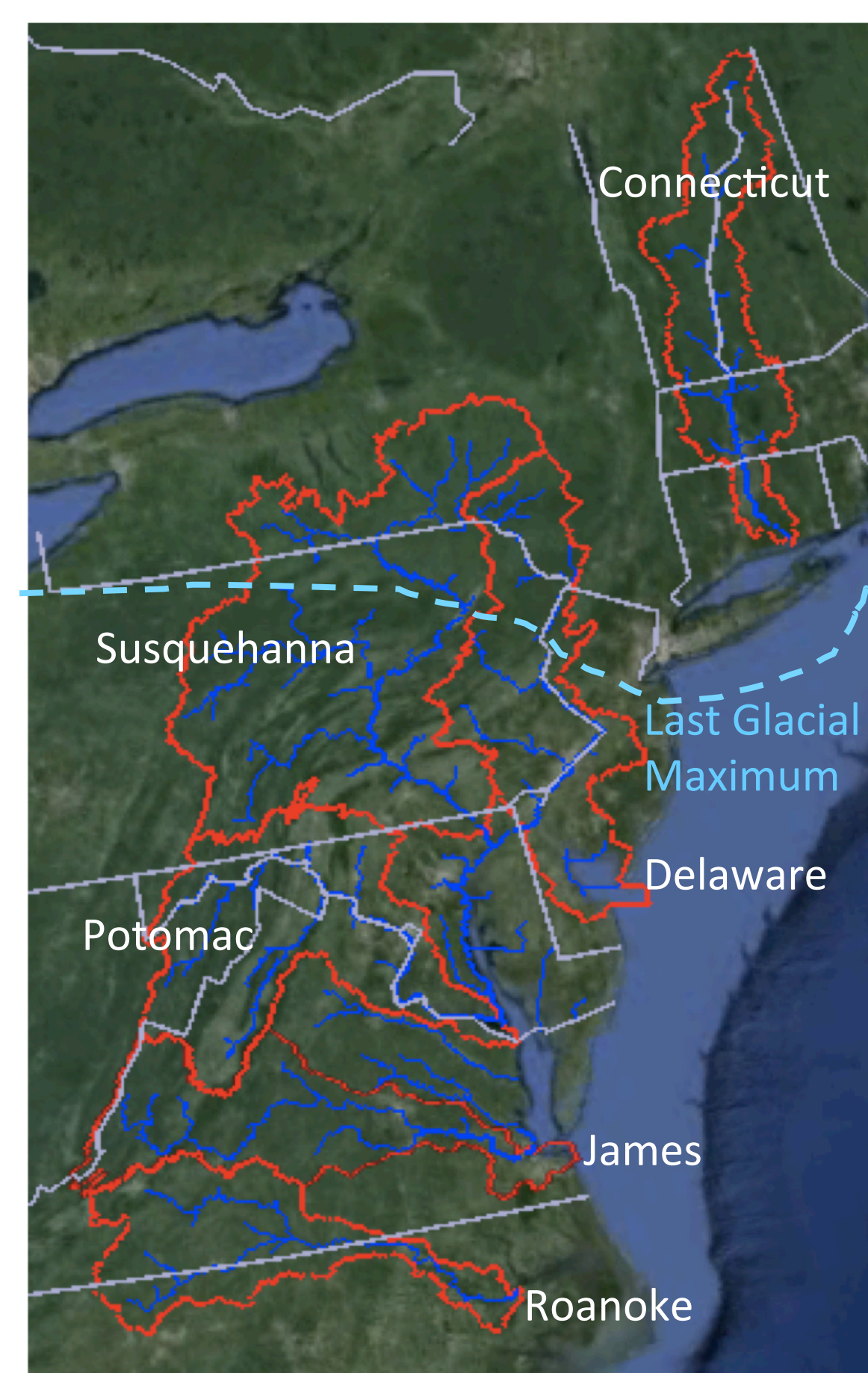
- Morphology of northern watersheds
- Hydro-climate of southern watersheds

High magnitude floods are larger in southern watersheds: Potomac, James, Roanoke

Implications

- Type of **flood risk** is tied to river valley morphology:
 - **Inundation** risk in southern rivers
 - **Erosion** risk in northern rivers
- Hydrologic Resiliency:
 - Northern river valleys *less resilient* to increasing frequency of large magnitude floods:
 - Narrow river valleys cannot accommodate large flood flows
 - Southern rivers are *more resilient to climate changes* **IF** floodplain morphology & functions are preserved: BUT urban development affects these functions

Study Region: Major Atlantic Slope River Basins



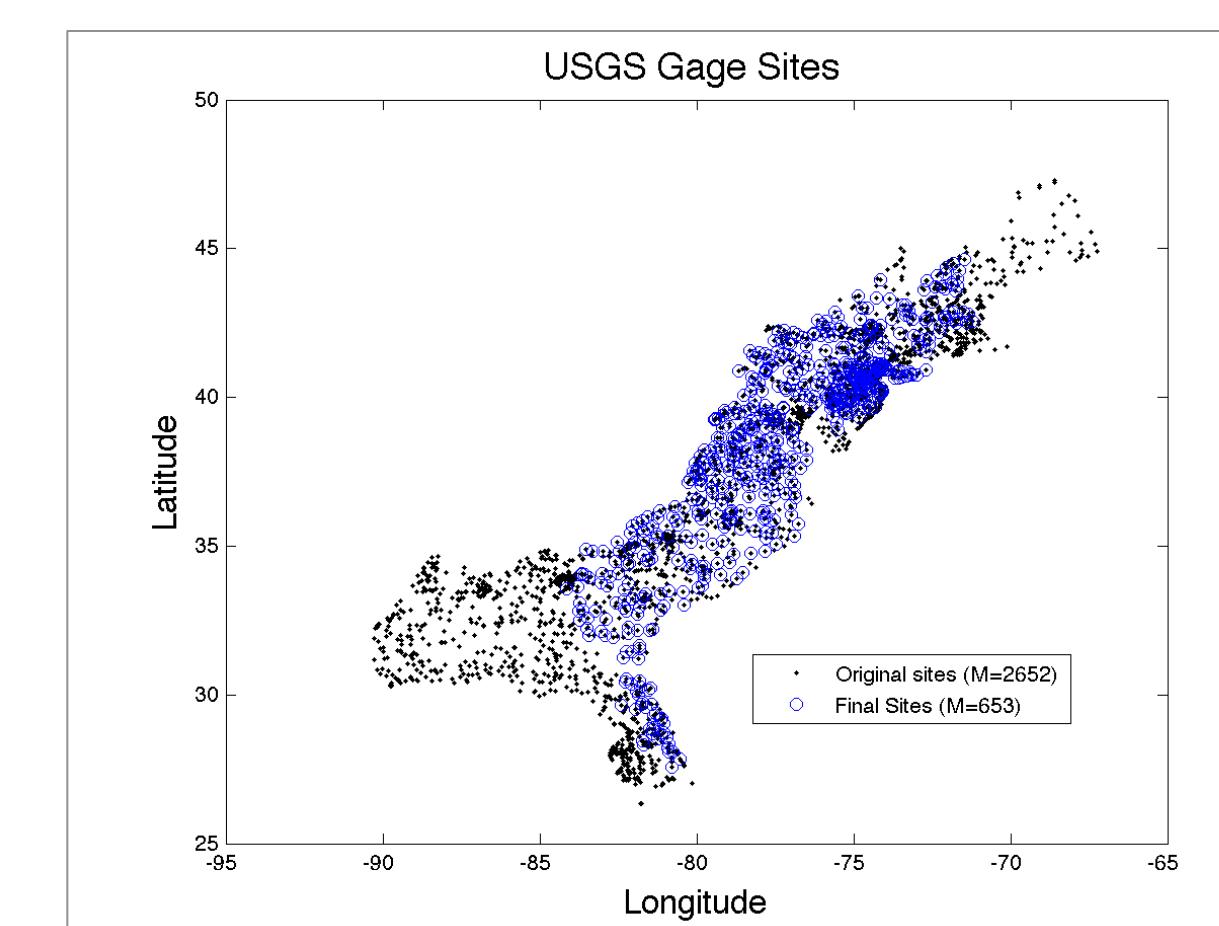
Geology:

- Northern region covered by Laurentide ice sheet at last glacial maximum
- Narrow bedrock valleys in north with no coastal plain
- Well-developed wide coastal floodplains in south

Hydro-climatology:

- Mixture of flood-producing precipitation events:
 - Tropical storms & hurricanes; summer convective storms; winter-spring extra-tropical storms
- Tropical storms dominate flood record in the south, while winter-spring storms dominate in the north

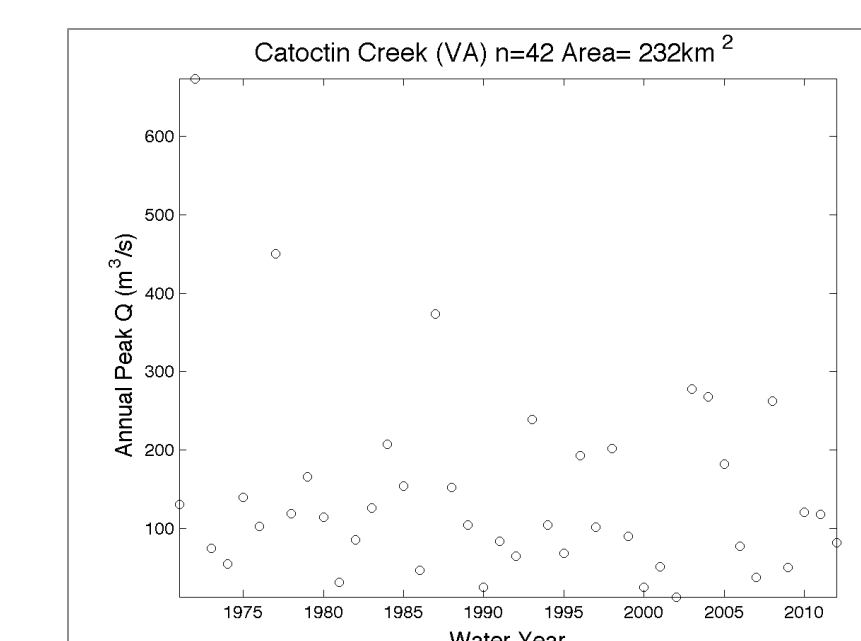
Methods



- Peak Flow Data from USGS Stream Gages: Annual Maximum Series
- Initial Criteria: all gages in New England, Mid-Atlantic & South Atlantic Hydrologic Regions
- Final criteria:
 - Min. 25 yrs of peak flow data
 - No major gaps in record
 - Flow not affected by regulation or channelization
 - No significant ($\alpha=1\%$) trends
 - Min. 10 gages in watershed
 - Watersheds drain to Atlantic

I. Flood Frequency Analysis

- Fit generalized extreme value distribution to annual maximum series at all gage sites
- Estimated Q_{RI} for range of recurrence intervals: 1.5 – 200 year floods



II. Spatial Scaling: $Q_{RI} = \alpha \log(A)^\theta$

- Linear regression of $\log(Q_{RI})$ on $\log(A)$ for each Q_{RI}
- Determined scaling exponent (θ) and coefficient (α) for each Q_{RI} in each river basin

